Environmental forcings of Paleogene Southern Ocean dinoflagellate biogeography

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Received 30 November 2009; revised 27 August 2010; accepted 1 October 2010; published 4 February 2011.

Despite warm polar climates and low meridional temperature gradients, a number of different high-latitude plankton assemblages were, to varying extents, dominated by endemic species during most of the Paleogene. To better understand the evolution of Paleogene plankton endemism in the high southern latitudes, we investigate the spatiotemporal distribution of the fossil remains of dinoflagellates, i.e., organic-walled cysts (dinocysts), and their response to changes in regional sea surface temperature (SST). We show that Paleocene and early Eocene (~65–50 Ma) Southern Ocean dinocyst assemblages were largely cosmopolitan in nature but that a distinct switch from cosmopolitan-dominated to endemic-dominated assemblages (the so-called “transantarctic flora”) occurred around the early-middle Eocene boundary (~50 Ma). The spatial distribution and relative abundance patterns of this transantarctic flora correspond well with surface water circulation patterns as reconstructed through general circulation model experiments throughout the Eocene. We quantitatively compare dinocyst assemblages with previously published TEX 86-based SST reconstructions through the early and middle Eocene from a key locality in the southwest Pacific Ocean, ODP Leg 189 Site 1172 on the East Tasman Plateau. We conclude that the middle Eocene onset of the proliferation of the transantarctic flora is not linearly correlated with regional SST records and that only after the transantarctic flora became fully established later in the middle Eocene, possibly triggered by large-scale changes in surface-ocean nutrient availability, were abundances of endemic dinocysts modulated by regional SST variations.


1. Introduction

[2] Early Paleogene temperatures and atmospheric CO2 concentrations were significantly higher than today [Pearson and Palmer, 2000; Sexton et al., 2006; Sluijs et al., 2006; Zachos et al., 2008; Bijl et al., 2009]. This long-term global warm period peaked during the late early Eocene (~52–50 Ma) [Zachos et al., 2001; Sexton et al., 2006], with high (>30°C) southern high-latitude SSTs and minimal latitudinal temperature gradients [e.g., Bijl et al., 2009]. The subsequent long-term global cooling through the middle and late Eocene [Sexton et al., 2006; Zachos et al., 2008; Bijl et al., 2009] culminated in the establishment of large Antarctic ice sheets around the time of the Eocene–Oligocene boundary (~34 Ma) [Miller et al., 1987; DeConto and Pollard, 2003a, 2003b; Coxall et al., 2005; Barker et al., 2007].

[3] The warm climates and low latitudinal temperature gradients prior to the onset of major Antarctic glaciation facilitated the poleward migration of many biota [Adams et al., 1990; Sluijs et al., 2006; Eberle et al., 2009]. Surprisingly, however, Southern Ocean fossil assemblages are distinctly endemic during long intervals of the early Paleogene and become increasingly endemic over time; besides mollusks [Zinsmeister, 1979], this phenomenon has been observed in numerous microfossil groups such as dinocysts [Wrenn and Beckman, 1982; Wrenn and Hart, 1988], radiolaria [Lazarus and Caulet, 1993; Lazarus et al., 2008], calcareous nanoplanckton and planktonic foraminifera [Nelson and Cooke, 2001; Villa et al., 2008], and diatoms [Harwood, 1991; Stickley et al., 2004b]. Dinocysts are
abundant in marginal marine sediments, notably at high latitudes. The early Paleogene Antarctic endemic dinocyst community is taxonomically well documented from outcrops and sediment cores around the Antarctic margin, and it is often referred to as the “transantarctic flora” [cf. Wrenn and Beckman, 1982]. However, the evolution and driving forces behind this Paleogene Southern Ocean endemism are largely unknown, although SST is presumed to have been involved [Huber et al., 2004]. So far, a lack of reliable SST reconstructions has hampered the evaluation of this hypothesis, and it is only through recently published geochemical data [Bijl et al., 2009] that such a test has become possible. [4] In this study, we aim to elucidate the role of oceanic surface circulation, SST and other environmental changes as a potential selective mechanism for Eocene Antarctic endemism among dinoflagellates. First, we provide an overview of Paleogene dinocyst paleobiogeography, for which we integrate newly generated palynological data with those previously published. Additionally, we statistically examine the relation of SST (using the TEX86 paleothermometry of Bijl et al. [2009]) and ocean-wide productivity (using the thallium isotope records of Nielsen et al. [2009]) to dinocyst assemblage changes in the southwest Pacific Ocean (Ocean Drilling Program Site 1172).

2. Material and Methods

2.1. Southern Ocean Sites

[5] We generated palynological data on sediments recovered at Ocean Drilling Program (ODP) and Deep Sea Drilling Project (DSDP) Sites 269, 277, 511, 512, 696, 739, 748 and 1090, as well as from the Brown’s Creek section in southeast Australia (Figure 1). Higher-resolution time series data were obtained from ODP Sites 1172 and 1171, and from outcrop samples from Blanche Point along the Southern Australian Margin. Furthermore, we reviewed published information on Paleogene (∼65–25 Ma) Southern Ocean dinocysts from Integrated Ocean Drilling Program (IODP), DSDP and ODP sites, the “Cape Roberts Project” (CRP 1, 2a and 3) and “Cenozoic Investigations of the Ross Sea 1” (CIROS-1) drilling expeditions. In addition, studies...
of outcrops that flank the Southern Ocean were considered from the southern Australian margin, McMurdo Sound, New Zealand, Tierra del Fuego, Santa Cruz Province, Brunswick Peninsula, and Seymour Island. An inventory of all sites and outcrops considered (including references) is listed in auxiliary material Table S1.\(^1\) An overview on the present-day locations of these sites is provided in Figure 1.

\(\text{[6]}\) The age control for these records is based on calcareous nannoplankton stratigraphy and, to a lesser extent, on diatom and/or palynomorph stratigraphies, radiometric dating, and magnetostratigraphy. The approximate tectonic configurations of the continents used in Figures 2 and 3 were reconstructed using the Ocean Drilling Stratigraphic Network (ODSN) database: http://www.odsn.de/ [after Hay et al., 1999].

\(\text{[7]}\) The semicontinuous succession recovered at ODP Site 1172 has been magnetostratigraphic calibrated [Stickley et al., 2004a] and extends from the Quaternary to the Maastrichtian (latest Cretaceous). Notably, it represents a reference section for Paleogene Antarctic endemism. ODP Site 1172 is located on the submerged western side of the East Tasman Plateau (ETP) at 45°S and a water depth of 2,620 m (Figure 1; see Exon et al. [2004b] for details). During the early Paleogene, the ETP was located at 65°S, much closer to Antarctica than it is today [Exon et al., 2004b]. Recovery at Hole 1172A includes Paleogene sediments down to the mid middle Eocene, while Hole 1172D extends the record from the mid middle Eocene to the Maastrichtian. Calcite concentrations increase from ∼40 Ma onward, which has been linked to a relative sea level rise following basin subsidence [Röhl et al., 2004]. Paleocene and Eocene sediments were recovered from between 350 and 700 m below seafloor (mbsf) [Exon et al., 2003; Stickley et al., 2004a]. This part of the section consists of brownish to greenish gray, bioturbated siltstones, that are poor in calcium carbonate, but yield abundant organic matter, including dinocysts [Exon et al., 2004b].

\(\text{2.2. Southern Ocean Surface-Current Configurations}\)

\(\text{[8]}\) According to modeling experiments and biogeographical data, in the early Paleogene the region was under the influence of the Antarctic-derived, northward flowing Tasman Current [Huber et al., 2004]. This current was part of a large gyre system in the South Pacific Ocean, the Proto-

\(\text{\textsuperscript{1}Auxiliary materials are available in the HTML: doi:10.1029/2009PA001905.}\)

**Figure 2.** Circum-Antarctic geographical distribution maps showing the spatiotemporal evolution of dinocyst endemism for seven different time slices. (a) Thanetian-Ypresian boundary (56–54 Ma), (b) mid-Ypresian (53–51 Ma), (c) Ypresian-Lutetian boundary (50–48 Ma), (d) early Lutetian (47–45 Ma), (e) late Lutetian (44–42 Ma), (f) Lutetian-Bartonian boundary (41–39 Ma), and (g) Bartonian-Priabonian boundary (38–36 Ma). Maps were derived from the Ocean Drilling Stratigraphic Network (ODSN). Black areas indicate (continental) blocks that are mostly subaerial. Note that several blocks shown in black were partly submerged during the Paleogene (e.g., Ross Sea, southern Australian margin, and parts of Argentina). Shaded areas indicate mostly submerged continental blocks [e.g., Brown et al., 2006]. For locality codes, see Table S1.
Ross Gyre [Stickley et al., 2004b]. The tectonic positioning of the continents that prescribe this current system was stable at least until ∼35.5 Ma, when the northward movement of Australia from Antarctica deepened the Tasman Gateway and much later allowed for the Antarctic Circumpolar Current [Stickley et al., 2004b].

2.3. Palynological Processing and Taxonomy

[9] Processing sediment samples for palynology followed the standardized methods described, e.g., by Sluijs et al. [2003]. Briefly, this involves treatment of dried sediment samples with ∼30% HCl and ∼38% HF. No bleaching and heavy-liquid separation was applied. A 15 μm nylon mesh was used for sieving. Residues were mounted on slides for microscope analysis. Approximately 200 dinocysts were counted per sample and identified to the species level at 500× magnification. An inventory of dinocyst species encountered in this study is listed in auxiliary material Table S2; dinocyst nomenclature and taxonomy, unless stated otherwise, are based on the work of Fensome and Williams [2004].

2.4. Dinocysts: Endemic Versus Nonendemic Taxa

[10] Dinocysts are the remains of unicellular eukaryotic plankton (dinoflagellates) that dwell as motile stages in surface oceans [e.g., Fensome et al., 1993]. Dinoflagellates occur in virtually all marine settings, but are most abundant and diverse in shelfal environments [e.g., Dale, 1996; Pross and Schmiedl, 2002; Pross and Brinkhuis, 2005; Sluijs et al., 2005]. The fossil record of dinoflagellates is predominantly based on their resistant organic-walled cysts (dinocysts). The dinocyst record, in turn, is particularly good in shelfal environments where relatively high sedimentation rates lead to a reduced oxygen exposure of the cysts. In contrast, deeper environments with well-ventilated bottom waters exhibit generally lower sedimentation rates and high oxygen concentrations, and in these settings dinocysts oxidize and do not preserve.

[11] Dinocyst species are considered to be “endemic to the Southern Ocean” if their empirically established distribution patterns are restricted to this region (i.e., latitudes south of 45°S). The Eocene endemic dinocysts in the Southern

Figure 2. (continued)
Ocean were originally referred to as the “transantarctic flora,” a name derived from their presumed affinity to the inferred Trans-Antarctic Seaway [Wrenn and Beckman, 1982]. Later studies reported elements of the transantarctic flora from other Southern Ocean sectors [Wrenn and Hart, 1988; Mao and Mohr, 1995; Levy and Harwood, 2000; Brinkhuis et al., 2003; Macphail and Truswell, 2004; Clowes and Wilson, 2006; Sluijs et al., 2009]. Some of these reports also described new species, but did not place those within the transantarctic flora. Other species occur in high abundances only in both north and south polar regions, and may therefore be placed in a “bipolar” group [e.g., Wrenn and Hart, 1988]. For this study, we include the bipolar species in our endemic dinocysts category herein after. Moreover, some taxa that are allegedly endemic to the Southern Ocean are morphologically similar to some cosmopolitan forms. For example, the supposedly endemic species Dracodinium waipawaense [Wilson, 1967] and the Northern Hemisphere D. varielongitudum [Williams and Downie, 1966] have the same stratigraphic range and are taxonomically identical, if one allows for subtle morphological variations in shape, size and number of processes (spines) on the cysts. Hence, it can be debated whether species such as D. waipawaense are truly endemic or whether the slight morphological differences are the result of ecophenotypic variability within a single dinocyst species. Because of these uncertainties, we did not consider such species to be part of the transantarctic flora. On the other hand, the Eocene species Enneadocysta dictyostila [Menéndez, 1965; Stover and Williams, 1995], emend.

Figure 3. Generalized middle to late Eocene (47–36 Ma) dinocyst geographical distribution map overlain by the ocean circulation pattern inferred from the GCM experiments of Huber et al. [2004]. Maps derived from the Ocean Drilling Stratigraphic Network (ODSN). Shaded areas indicate mostly submerged continental blocks [Brown et al., 2006]. Abbreviations: TA-SW, Trans-Antarctic Seaway (hypothetical [see Wrenn and Beckman, 1982]); TSA-SW, Trans-South American Seaway (hypothetical [see Kohn et al., 2004; Lagabrielle et al., 2009]); EAC, East Australian Current; p-LC, Proto-Leeuwin Current; p-RG, proto-Ross Gyre; TC, Tasman Current [Huber et al., 2004].
Table 1. Overview of Dinocysts Herein Assigned to the Transantarctic Flora and the Bipolar Species

<table>
<thead>
<tr>
<th>New Endemic Species</th>
<th>Species Now Assigned to Transantarctic Flora</th>
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<tbody>
<tr>
<td>Described prior to Wrenn and Hart [1988]</td>
<td>Alterbidinium distinctum</td>
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<td></td>
<td>Arachnodochnium antarcticum</td>
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<td></td>
<td>Deflandrea antarctica</td>
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<td></td>
<td>Deflandrea cygniformis</td>
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<td>Deflandrea granulata</td>
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<td>Deflandrea oehsseldensis</td>
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<td></td>
<td>Spindinium macmurdense</td>
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<td></td>
<td>Vozzhennikovia apertura</td>
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<tr>
<td></td>
<td>Wilsonidium echinosuturatum</td>
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<tr>
<td>Described by Wrenn and Hart [1988]</td>
<td>Spindinium colemani</td>
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<td></td>
<td>Octodonim askiniae</td>
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<td></td>
<td>Hystrichosphaeridium trueswelliae: now nonendemic [Sluijs and Brinkhuis, 2009].</td>
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<td></td>
<td>Impletothesphaeridium clavus</td>
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<td></td>
<td>Deflandrea webbii</td>
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<td></td>
<td>Spindinium luciae</td>
</tr>
<tr>
<td>Described by Goodman and Ford [1983]</td>
<td>Phthanoperidinium sp. A</td>
</tr>
<tr>
<td>Described by Stover and Williams [1995]</td>
<td>Gipslandia extensa</td>
</tr>
<tr>
<td>Described by Max and Mohr [1995]</td>
<td>Phthanoperidinium antarcticum</td>
</tr>
<tr>
<td>Described by Stover and Williams [1995]</td>
<td>(now) Enneadocysta dictyostila</td>
</tr>
<tr>
<td>Described by Levy and Harwood [2000]</td>
<td>Vozzhennikovia netrona</td>
</tr>
<tr>
<td>Described by Brinkhuis et al. [2003]</td>
<td>Deflandrea sp. A</td>
</tr>
<tr>
<td>Described by Clowes and Wilson [2006]</td>
<td>Corrudinium regulare</td>
</tr>
<tr>
<td>Described by Sluijs et al. [2009]</td>
<td>Vozzhennikovia stickleyae</td>
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<td></td>
<td>Vozzhennikovia roehliae</td>
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<td></td>
<td>Spinidinium schellenbergii</td>
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<td>Moria zachosi</td>
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<tr>
<th>Bipolar Species</th>
<th>Species Now Considered to Be Bipolar</th>
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<tr>
<td>Assigned bipolar by Wrenn and Hart [1988]</td>
<td>(now) Phthanoperidinium</td>
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<td>Assigned bipolar by Warnaar [2006]</td>
<td>stockmansi group</td>
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<tr>
<td>Assigned bipolar by Wrenn and Hart [1988]</td>
<td>Spindinium macmurdense</td>
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3. Results

3.1. Dinocyst Biogeography Through Time

[13] Distribution maps of the compiled circum-Antarctic dinocyst records for seven time slices spanning ~56 Ma to 36 Ma are presented in Figures 2a–2g. The distribution patterns for the two oldest time slices (Figure 2a, 56–54 Ma; Figure 2b, 53–51 Ma) have limited numbers of dinocyst-bearing sites restricted to the south Pacific Ocean and Drake Passage. This limited information shows that endemics are present, but do not numerically dominate the dinocyst assemblages, not even on the Antarctic Margin. During the early middle Eocene, endemic taxa are dominant in the areas east of the Tasmanian Gateway and Drake Passage (Figure 2c, 50–48 Ma). From the middle Eocene onward, endemic dinoflagellates consistently and strongly dominate the assemblages along the western boundaries of the Drake Passage and the Tasman Gateway as well as at Prydz Bay (Figures 2d, 2e, and 2f, 47 Ma to 39 Ma), at least until the mid to late Eocene (Figure 2g, 38–36 Ma). Many undoubtedly Oligocene and younger records in the southwest Pacific (not presented here; e.g., ODP Leg 28 [Kemp, 1975]) and Atlantic [Guerstein et al., 2008] yield some endemic dinocysts. These are most likely reworked from Eocene sediments as a result of Oligocene glacial activity. We have compiled a composite of the generalized distribution of endemic dinocysts during the middle and late Eocene (47–36 Ma) and note that the sectors characterized by dominant endemism are restricted to the Antarctic margins and the eastern sides of the Tasman and Drake conduits (Figure 3). The Australo-Antarctic Gulf area is characterized...
by overall low abundances of endemics throughout the Paleogene (Figure 3).

3.2. Statistical Analysis

With our DCCA analyses, four canonical axes are established for the four environmental variables. The first axis explains 29% of the dinocyst data, the second axis another 4%. Figure 4 shows the dinocyst scores for the first two axes with the environmental vectors. The SST vector points to the northeast, whereas the other vectors (i.e., thallium isotopes, % peridinioid dinocysts and % endemic dinocysts) are approximately perpendicular to the temperature vectors.

In Figure 4, the dinocyst groups are divided into endemic (blue triangles) and nonendemic (red circles). The endemic taxa are clearly separated from the nonendemic ones, with high scores on both axes for nonendemic and low scores for endemic dinocysts. The few exceptions include Membranophoridium perforatum (Mper) and Hystrichosphaeridium truswelliae (Htru). The latter has traditionally been interpreted as being endemic to the Southern Ocean. However, recently H. truswelliae was recorded in Paleocene-Eocene strata from New Jersey (northeast North America), clearly disputing its endemic signature [Sluijs and Brinkhuis, 2009]. ‘Outliers’ within the nonendemic taxa include the cosmopolitan Enneadocysta species (Enne), Phthanoperidinium comatum (Pcom) and Selenopemphix spp. (Sele). Comparing the scores of the first two axes to the environmental variables, we note that the SST variations correspond with the sample scores of the second axis. This axis explains only 4% of the dinocyst variation (Figure 5). Patterns in the sample scores of the first axis correspond with the abundance of endemic dinocysts (Figure 5). Thus, while spatial gradients in % endemic dinocysts appear to be a good indicator of the location of ocean currents, SST appears to have had little effect on the biogeography of endemic dinocysts over the complete record.

However, during the interval of long-term cooling which began after the Early Eocene Climate Optimum (EECO) [Bijl et al., 2009], abundance changes in endemic dinocysts do now covary with SST. For instance, during the Middle Eocene Climatic Optimum (MEO), a short-lived warming event at ~40 Ma [Bohaty et al., 2009], endemic dinocyst abundances fall again to below 50% (Figure 5). After the MECO, the cooling trend continues and endemic dinocysts redominate the assemblages (Figure 5). To test the relation between dinocyst endemism and SST in the post-EECO interval of endemic dinocyst dominance, we performed another DCCA analysis involving only those data...
within the middle Eocene interval. In this interval, the variation explained by SST is ∼20% (Figure 6).

4. Discussion

4.1. Comparison With Other Fossil Groups

While endemism is recorded within several fossil groups in the Eocene, a direct comparison with the dinocyst record of the southwest Pacific is hampered by a paucity of corresponding records around the Antarctic margin older than ∼45 Ma. For instance, although diatomaceous (biosiliceous) sediments of Early Cretaceous age and K-T boundary age occur in the western Antarctic margin [Harwood, 1988; Gersonde and Harwood, 1990; Harwood and Gersonde, 1990], these appear to be a rare exception since most known biosiliceous records in Antarctic marginal (shelf) areas are mid Eocene in age (i.e., ODP Site 1172 [Stickley et al., 2004a]) or commonly much younger [e.g., Lazarus et al., 2008]. Calcareous microfossil data from the southern high latitudes predating 45 Ma are similarly scarce. Nevertheless, the few existing records of microfossils other than dinocysts from that time indeed show an increase in endemism beginning in the mid middle Eocene, as has been shown for radiolarians [Lazarus et al., 2008] and calcareous nannoplankton [Villa et al., 2008]. Calcareous nannoplankton also exhibit a decrease in endemism during the MECO [Villa et al., 2008].

4.2. Potential Driving Forces: SST, Endemism, and Statistics

The establishment of dominant Antarctic endemic dinocyst assemblages characterizes the early middle Eocene transition (∼50 Ma). However, SSTs did not decrease abruptly at that time (Figure 5). This implies that the increase in abundance of endemic dinocysts was not linearly related to SST. In contrast, during the middle Eocene, abundances of endemic dinocysts are modulated by SST changes (Figure 6). What, then, triggered endemic dinocysts
to proliferate shortly after ~50 Ma? As stated above, the peak warmth of the EECO at ~50 Ma also marks the transition into the onset of high-latitude and global cooling that, despite an interruption by at least one short warming interval (i.e., the MECO, at ~40 Ma; Figure 5), culminated in the onset of Antarctic glaciation during the Eocene-Oligocene transition [e.g., Zachos et al., 1992]. It may be possible that the endemic dinoflagellate taxa, when they appeared, were capable to sustain in warm subtropical conditions during the EECO, but subsequently evolved to be more specialized to cooler surface waters during middle Eocene cooling. The sudden shift to dominant endemic dinocysts suggests the crossing of an environmental threshold at which cosmopolitan taxa were outcompeted by the endemic taxa. Because the DCCA statistical method identifies linear correspondence between assemblages and environmental data, this threshold may still have been SST.

[19] The mechanism described above may provide an explanation for the proliferation of endemic dinocysts just after 50 Ma, but it fails to explain the presence, in low abundance, of endemic dinocysts during Paleocene and early Eocene times. The most abundant endemic dinocyst groups, i.e., Deflandrea antarctica, Ph thannerperidinium stockmansii group and Vozzhennikovia/Spinidinium spp., are already present in the Paleocene. Apparently, Paleocene—early Eocene surface waters were not as favorable for endemic species as they were in the middle Eocene. Hence, along with SST changes, other physicochemical parameters of the surface waters must have changed between the Paleocene and middle Eocene that stimulated Southern Ocean endemism to proliferate just after the early middle Eocene boundary. A possible explanation may lie in the diatom record at ODP Site 1172, where opal-A is preserved from approximately 47 Ma [Stickley et al., 2004a, 2004b] (or 45 Ma [Bijl et al., 2009]) onward (Figure 7). In the middle Eocene, global cooling enhanced latitudinal gradients [Bijl et al., 2009]. This cooling stimulated upwelling, bringing silica-enriched waters that characterized Eocene oceans [e.g., McGowran, 1989], particularly the Pacific [e.g., Moore et al., 2008], to the surface for utilization by the diatoms. This change apparently brought about significant diatom production and preservation to the region at ODP Site 1172, and coincides with a global evolutionary turnover event within the diatoms and globally enhanced preservation of opal-A associated with global cooling [Barron and Baldauf, 1989; McGowran, 1989; Baldauf and Barron, 1990]. Increased weathering may have led to an additional nutrient supply to the oceans including both Si and the diatom bio-limiting nutrient Fe [Boyd et al., 2000]. It appears plausible that the endemic dinocysts that proliferated during the earliest middle Eocene benefited from the same circulation changes and eutrophication phase that stimulated diatom productivity at that time. Further evidence for changes in ocean-wide surface water fertility comes from thallium isotopes measured on ferromanganese crusts from the Pacific Ocean (Figure 7) [Nielsen et al., 2009]. A shift in these isotopes together with a coeval shift in sulfur isotopes suggests a major ocean-wide increase in marine organic carbon export from 55 to 45 Ma. At the same time, deep-water carbon isotopic values between oceanic basins start to diverge, suggesting (among other possibilities) an increase in the biological pump leading to increased water mass aging gradients (Figure 7) [Sexton et al., 2006]. The fertilization of the oceanic basins is, as suggested

Figure 6. (left) Middle Eocene TEX86-based SST curve from ODP Site 1172 [Bijl et al., 2009] overlain by the first axis scores of the middle Eocene DCCA experiment (in red; 20% of the dinocyst variation). The second axis’s scores correspond to (right) the % endemic dinocyst curve but explain only 4.7% of the variation in the dinocyst assemblage.
5. Conclusions

[20] New data and reviews of previous reports allow a considerably improved spatiotemporal reconstruction of Paleogene circum-Antarctic dinocyst distribution patterns. Dinocysts can be divided into Antarctic endemic, bipolar, middle- to low-latitude, and cosmopolitan groups. Their spatial distribution patterns match ocean current patterns generated through GCM experiments [Huber et al., 2004] during the middle to late Eocene. Despite open connections between the Southern Ocean and adjacent midlatitude oceans, [Exon et al., 2004a], circum-Antarctic dinocyst assemblages switched from cosmopolitan and low-latitude-derived dinoflagellate cyst assemblage to an endemic-dominated assemblage in the early middle Eocene. By numerically evaluating the relation between Southern Ocean dinocyst assemblages and the SST evolution at ODP Site 1172, we identify that SST was not the primary driving factor in the establishment of dominantly endemic dinocyst assemblages. We conjecture that ocean-wide surface water ocean fertility [Sexton et al., 2006; Nielsen et al., 2009] brought about at post-EECO cooling [Sexton et al., 2006; Bijl et al., 2009], helped to stimulate diatom production and ‘switch on’ opal preservation in the area and caused a rapid reorganization of high-latitude dinocyst assemblages. Once the endemic community became the dominant component of the dinocyst assemblages in the early middle Eocene southwest Pacific, the influence of SST variations on the dinocyst assemblage was stronger than during the Paleocene and early Eocene.

[21] Acknowledgments. This research used samples and data provided by the Deep Sea Drilling Project (DSDP) and the Ocean Drilling
Program (ODP). DSDP and ODP were sponsored by the U.S. National Science Foundation (ODP was also sponsored by other participating countries) under the management of Joint Oceanographic Institutions (JOI) Inc. The Shipboard Scientific Party of IODP Expedition Leg 318 is thanked for providing data. This project was funded by the LPP foundation to P. K. Bijl. J. P. Houben acknowledges support by the Biodiversity and Climate Research Center of the Hessian initiative for scientific and economic excellence (LOEWE). C. E. Sticklekeys acknowledges the Research Council of Norway. R. Guerstein acknowledges the Agencia Nacional de Promoción Científica y Tecnológica–PICT 26057. A. Sluijs thanks the Netherlands Organization for Scientific Research (Veni grant 863.07.001). M. Huber acknowledges continuing support from NSF 0927946-ATM and NSF PC2C 0900282- OCE. M. Huber is also grateful to the New Zealand GNS Science for providing an excellent work environment while he was visiting. A. J. P. Houben and H. Brinkhuis are grateful to Statoil for financial support. We thank Giuseppe Cortese, Erica Crouch (both GNS Science, New Zealand), Oliver Hein (Utrecht University), and Martin Pearce (Statoil) for constructive discussions. We thank Jan van Tongeren and Natasja Walters for technical support. Constructive reviews by Philip Sexton and an anonymous reviewer greatly an earlier version of the manuscript.

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